




# Global tree intrinsic water use efficiency is enhanced by increased atmospheric CO<sub>2</sub> and modulated by climate and plant functional types

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We conducted a meta-analysis of carbon and oxygen isotopes from tree ring chronologies representing 34 species across 10 biomes to better understand the environmental drivers and physiological mechanisms leading to historical changes in tree intrinsic water use efficiency (iWUE), or the ratio of net photosynthesis ( $A_{\text{net}}$ ) to stomatal conductance ( $g_s$ ), over the last century. We show a ~40% increase in tree iWUE globally since 1901, coinciding with a ~34% increase in atmospheric CO<sub>2</sub> ( $C_a$ ), although mean iWUE, and the rates of increase, varied across biomes and leaf and wood functional types. While  $C_a$  was a dominant environmental driver of iWUE, the effects of increasing  $C_a$  were modulated either positively or negatively by climate, including vapor pressure deficit (VPD), temperature, and precipitation, and by leaf and wood functional types. A dual carbon–oxygen isotope approach revealed that increases in  $A_{\text{net}}$  dominated the observed increased iWUE in ~83% of examined cases, supporting recent reports of global increases in  $A_{\text{net}}$ , whereas reductions in  $g_s$  occurred in the remaining ~17%. This meta-analysis provides a strong process-based framework for predicting changes in tree carbon gain and water loss across biomes and across wood and leaf functional types, and the interactions between  $C_a$  and other environmental factors have important implications for the coupled carbon–hydrologic cycles under future climate. Our results furthermore challenge the idea of widespread reductions in  $g_s$  as the major driver of increasing tree iWUE and will better inform Earth system models regarding the role of trees in the global carbon and water cycles.

tree rings | stable isotopes | carbon | oxygen | intrinsic water use efficiency

How terrestrial plants respond to more frequent, and often prolonged, environmental stressors will have profound impacts on, and feedbacks to, the Earth–climate system at regional to continental scales (1, 2). Central to these feedbacks are plant stomata, microscopic pores on the leaves of plants that act as a control valve over the fluxes of carbon dioxide ( $C_a$ ) into the leaf during photosynthesis and water vapor ( $H_2O$ ) out of the leaf during transpiration. Importantly, changes in stomatal aperture do not affect the fluxes of  $C_a$  and  $H_2O$  equally, as the sum of resistances for the diffusion of  $C_a$  from the atmosphere to mesophyll cells where Rubisco is located are much greater than those for  $H_2O$  from the surface of leaf mesophyll cells to the atmosphere (3). Indeed, as stomatal aperture changes, so does water use efficiency (WUE), or the ratio of  $C_a$  uptake to  $H_2O$  released from the leaf to canopy scale (4). Consequently, understanding the environmental factors driving changes in leaf physiology is of paramount concern in the context of climate change as small changes in tree WUE can have major effects on the carbon and hydrologic cycles over large geographical areas (1, 5).

Approaches using tree ring carbon isotopes (6–9), eddy-flux measurements (4, 9, 10), atmospheric carbon isotope composition analysis (11), and Earth system modeling techniques (9–13) have shown trends of recently increasing WUE. These increases can occur by stimulation of leaf photosynthetic rates ( $A_{\text{net}}$ ) (14, 15), reduced stomatal conductance to water ( $g_s$ ) (14, 15), or some combination of the two. A fundamental physiological response

found in numerous  $C_a$  enrichment experiments is that WUE of many plants is improved as a result of increasing  $C_a$  stimulating photosynthesis and causing partial stomatal closure (14, 16). However, environmental factors distinct from  $C_a$ , such as vapor pressure deficit (VPD), precipitation, and temperature, have independent effects on  $A_{\text{net}}$  and  $g_s$  and, therefore, may modulate the response of WUE to rising  $C_a$ , especially across functionally distinct plant groups with differences in wood anatomy (9) and leaf morphology (17). Despite this, few studies have thoroughly examined the effects of multiple environmental factors over controls of WUE, and even fewer have considered the underlying component parts,  $A_{\text{net}}$  and  $g_s$  (9, 13, 17–20). This has, in part, been due to the complexity of partitioning  $H_2O$  gas fluxes at ecosystem scales (21), in addition to the difficulty in attributing changes in isotopically derived intrinsic water use efficiency (iWUE) (the ratio of  $A_{\text{net}}$  to stomatal conductance to water,  $g_s$ ) to  $A_{\text{net}}$  or  $g_s$  without the accompaniment of physiological measurements (18, 19).

A promising technique couples carbon isotopically derived estimates of iWUE with oxygen isotope leaf water enrichment above source water ( $\Delta^{18}O_{\text{lw}}$ ; derived from tree ring  $\delta^{18}O$  and source water  $\delta^{18}O$  assumed to be  $\sim\delta^{18}O_{\text{precipitation}}$ ) to provide a qualitative attribution of changes in iWUE to underlying  $A_{\text{net}}$  and  $g_s$  (9, 20, 22–29). As  $\Delta^{18}O_{\text{lw}}$  is inversely related to  $g_s$  (26–30), if increases in iWUE were due to increases in  $A_{\text{net}}$ , then  $\Delta^{18}O_{\text{lw}}$  should be constant or decrease with iWUE. However, if increases in iWUE were due to decreases in  $g_s$ , or a combination of a decrease in  $g_s$  and an increase in  $A_{\text{net}}$ ,  $\Delta^{18}O_{\text{lw}}$  would increase

## Significance

Changes in tree physiology driven by environmental change can alter the balance of forest ecosystem carbon and water fluxes. We performed a meta-analysis of published tree ring literature, comprising 36 different species across 84 sites globally, to show stimulated leaf photosynthesis, not reduced stomatal conductance, is primarily responsible for recently increasing tree intrinsic water use efficiency (iWUE), which integrates the balance between carbon and water fluxes. Furthermore, we show trends in tree iWUE are similar in magnitude to the increase in atmospheric CO<sub>2</sub> over the 20th century and that climate interacts with CO<sub>2</sub> to modulate tree iWUE. These findings will help to guide efforts of refining the role of forests in process-based models under future environmental change.

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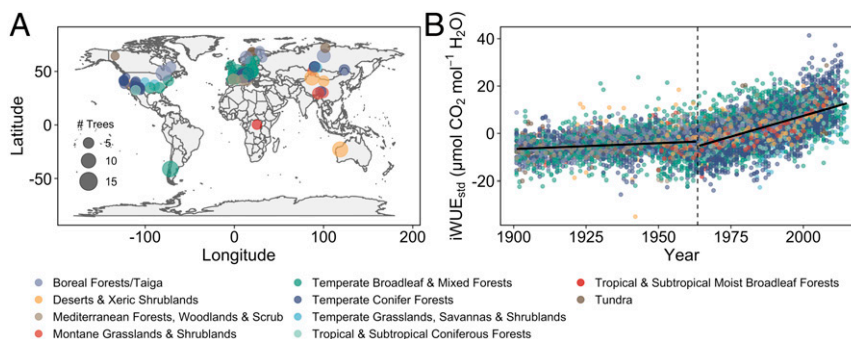
with iWUE (9, 25, 30). As such, there currently exists a wealth of previously untapped long-term records of tree physiological responses to environmental change within numerous dendrochronological studies from around the world, providing a historical view of how iWUE has changed globally over the last century. In this analysis, we 1) synthesize published data from tree ring carbon and oxygen isotope chronologies to examine global trends in tree ring-derived iWUE, 2) identify those environmental factors, and their interactions, that best explain multidecadal to centennial trends in iWUE, and 3) investigate the potential underlying changes in  $A_{\text{net}}$  and  $g_s$  through an analysis of coupled tree ring-derived iWUE and  $\Delta^{18}\text{O}_{\text{w}}$  over time (9, 20, 22, 23, 25).

Using 113 unique tree ring carbon and oxygen isotope chronologies comprising 36 different species across 84 sites globally (Fig. 1A and SI Appendix, Table S1), we show tree-level iWUE increased, on average, by ~40% ( $0.35\% \text{ y}^{-1}$ ,  $\text{iWUE} = 0.23 \cdot \text{y} - 369.16$ ) over the last century (1901–2015) (Fig. 1B). We statistically identified a breakpoint in the combined iWUE chronology at 1963, after which iWUE increased linearly at a rate of  $0.39 \pm 0.01 \mu\text{mol CO}_2\text{-mol}^{-1} \text{ H}_2\text{O}\cdot\text{y}^{-1}$  ( $1.67\% \text{ y}^{-1}$ ), or ~3.9 times faster than the previous 63 y ( $F = 207.14$ ,  $P < 0.0001$ ) (Fig. 1B and SI Appendix, Fig. S1). This change in the rate at which tree iWUE increased after 1963 coincided with a statistical breakpoint in  $C_a$  at 1969 where  $\text{CO}_2$  began increasing 4.1 times faster than between 1901 and 1969 (SI Appendix, Fig. S2) (31), and is similar to the trend in iWUE from a recent global synthesis by Adams et al. (32) who also showed a similar breakpoint in the 1960s. When considering individual chronologies, the increases in iWUE were widespread, with 93% (105 of 113 chronologies) of those examined having positive trends over 1901–2015 and 84% (95 of 113) of those examined having positive trends over 1963–2015 (SI Appendix, Table S1). These data for trees during the Anthropocene spanning 10 biomes in six continents that represent a spectrum of leaf and wood types reinforce reports of increasing iWUE in the United States (9), Europe (6, 8), and tropical forests (7), in addition to a recent compilation of global iWUE trends (32).

Globally, differences in vegetation physiognomy may have large impacts on iWUE, and despite much work examining plant physiological processes within biomes, there have been few large-scale comparisons of historical tree iWUE across biomes at a global scale (6, 8, 9, 12, 32). In this study, we found the rate of increase of iWUE from 1963 to 2015 differed among the 10 biomes represented ( $F = 7.21$ ,  $P < 0.001$ ) and ranged between  $0.101 \pm 0.07 \mu\text{mol CO}_2\text{-mol}^{-1} \text{ H}_2\text{O}\cdot\text{y}^{-1}$  for trees growing in deserts and xeric shrublands to  $0.611 \pm 0.07 \mu\text{mol CO}_2\text{-mol}^{-1} \text{ H}_2\text{O}\cdot\text{y}^{-1}$  for Mediterranean forests, woodlands, and scrub (SI Appendix, Fig. S3A and Table S2). Like deserts and xeric shrublands, the rate of

iWUE increase since 1963 for trees growing in the tundra and temperate grasslands, savannas, and shrublands was low, with a mean across the three biomes of  $0.131 \mu\text{mol CO}_2\text{-mol}^{-1} \text{ H}_2\text{O}\cdot\text{y}^{-1}$ . All other biomes exhibit mean rates of iWUE increase after 1963 greater than  $0.409 \mu\text{mol CO}_2\text{-mol}^{-1} \text{ H}_2\text{O}\cdot\text{y}^{-1}$  (SI Appendix, Table S2). Furthermore, mean iWUE over 1963–2015 ranged from a low of  $59.9 \mu\text{mol CO}_2\text{-mol}^{-1} \text{ H}_2\text{O}$  in tropical and subtropical moist broadleaf forests to a high of  $91.5 \mu\text{mol CO}_2\text{-mol}^{-1} \text{ H}_2\text{O}$  in temperate conifer forests (SI Appendix, Fig. S3A).

When considering changes in iWUE since 1963 with respect to wood anatomy, the rate of iWUE increase was ~21% higher for conifers relative to diffuse porous trees ( $P = 0.022$ ), but we found no difference in the rate of iWUE increase between conifer trees and ring porous trees ( $P = 0.61$ ) or between diffuse porous trees and ring porous trees ( $P = 0.27$ ) (SI Appendix, Fig. S4A). Our results are similar to climate-corrected iWUE trends from tree rings presented by Frank et al. (6) in European forests, but are in contrast to work by Saurer et al. (33) and Wang et al. (34) using tree rings and flux tower measurements, respectively, who showed broadleaf deciduous trees (i.e., ring porous and diffuse porous) in the Northern hemisphere having greater rates of increase in iWUE than conifer trees over the last century. It is important to note, however, that minor discrepancies among absolute iWUE values and their trends over time in these studies when compared to ours may have resulted from the formulation (i.e., our inclusion of a photorespiratory term) (35) and methodology differences (i.e., tree ring, flux tower) (36). We found no differences in the rate of iWUE increase since 1963 among trees with different leaf functional types ( $F = 1.05$ ,  $P = 0.37$ ) (SI Appendix, Fig. S4C). Additionally, we found mean iWUE during 1963–2015 was different among all wood types ( $F = 782.96$ ,  $P < 0.001$ ), being the highest in conifer species and lowest in ring porous species (SI Appendix, Fig. S4B), consistent with recent findings from 12 tree species at eight forested sites in the United States (9). Of important note, the patterns in mean iWUE for each wood type fall along a gradient of hydraulic conductivity and safety trade-offs (37). Conifer trees, which have relatively low hydraulic conductivity, but are more resistant to drought induced cavitation and embolism, had the highest mean iWUE, in contrast to ring porous trees, which have higher hydraulic conductivity, but are more vulnerable to hydraulic failure, which had lower mean iWUE (38–40). Mean iWUE was also different among trees with different leaf types, with needleleaf evergreen trees being the highest, followed by needleleaf deciduous trees, broadleaf evergreen trees, and finally broadleaf deciduous trees (SI Appendix, Fig. S4D).



**Fig. 1.** Tree locations from which chronologies of iWUE and  $\Delta^{18}\text{O}_{\text{w}}$  were developed (A) and group mean-centered iWUE by species within site over the period 1901–2015 (B). The color of data points in A and B correspond to the biome from which trees were growing, while the size of the circle in A corresponds to the number of trees used in the development of each carbon and oxygen isotope-derived chronology, respectively. The vertical dashed line in B occurs at the year 1963 where the rate of change in iWUE increases. The solid lines in B denote the average trend in iWUE for the period before (1901–1963) and after (1963–2015) the identified breakpoint. Data for iWUE trends and the number of trees per chronology are listed in SI Appendix, Table S1.

Changes in climate and  $C_a$  have strong effects on vegetation function from the leaf (14, 15) to global scale (13, 41, 42), yet how environmental change has influenced iWUE across large spatial scales over the last century is not fully resolved. Guerrieri et al. (9) and Frank et al. (6) recently showed increasing  $C_a$  and climate change have led to increased iWUE in tree species across the United States and Europe, but whether this response is conserved across biomes experiencing a much larger range in climate is still unknown. To address this knowledge gap, we used linear mixed effects (LME) models to examine the importance of environmental factors in driving tree iWUE. Across all species and study locations from 1963–2015, LME model results explained as much as ~89% of the variance in iWUE and indicated a strong, positive relationship with  $C_a$  and growing season vapor pressure deficit ( $VPD_{grw}$ ) and a negative relationship with growing season precipitation ( $PPT_{grw}$ ), although a large proportion of the total variance was attributed to differences among sites (Table 1 and *SI Appendix, Table S3*). Remaining unaccounted variance may include non-climatic factors, such as nitrogen availability or acidic air pollution, which are known to have important influences over iWUE (19, 43, 44), but these were not included in the studies from which we extracted isotopic data. When extending this analysis to changes in iWUE over the last 115 y (1901–2015), LME model main effects were remarkably consistent with data from 1963–2015, with the exception of growing season temperature ( $TMP_{grw}$ ) having a marginally positive effect on iWUE across the 115-y chronology (*SI Appendix, Table S4*), which suggests  $TMP_{grw}$  is becoming more important in recent years.

As a complement to our analysis using LME models to examine the effects of  $C_a$ ,  $PPT_{grw}$ ,  $TMP_{grw}$ , and  $VPD_{grw}$  on tree iWUE, we used hierarchical partitioning (HP), which alleviates potential problems that arise due to multicollinearity, to estimate the individual contribution of a given environmental parameter to tree iWUE. For the period 1901–2015, HP indicated that, of the 42% variance in iWUE explained by our model,  $C_a$  accounted for 59.6% ( $z = 1,375.5$ ,  $P < 0.05$ ), more than any other factor considered. We

found  $VPD_{grw}$ ,  $TMP_{grw}$ , and  $PPT_{grw}$  contributed 23.6% ( $z = 501.3$ ,  $P < 0.05$ ), 16.2% ( $z = 340.1$ ,  $P < 0.05$ ), and 0.6% ( $z = 12.6$ ,  $P < 0.05$ ), respectively. Of the 31% variance in iWUE explained in our model over the period 1963–2015, the influence of  $C_a$  declined to 45.2% ( $z = 455.6$ ,  $P < 0.05$ ), consistent with Adams et al. (32), who found a diminishing rate of response of iWUE to  $CO_2$  after 1966 (*SI Appendix, Fig. S5*). On the other hand, the influence of  $TMP_{grw}$  increased to 28.0% ( $z = 357.2$ ,  $P < 0.05$ ), whereas the contributions of  $VPD_{grw}$  and  $PPT_{grw}$  were similar to the 1963–2015 chronology with 26.1% ( $z = 269.7$ ,  $P < 0.05$ ) and 0.7% ( $z = 6.4$ ,  $P < 0.05$ ), respectively.

Increasing  $C_a$  may affect both stomatal conductance and photosynthesis (14, 15). VPD, or dryness of the atmosphere, regulates stomatal conductance (45), and therefore, both photosynthesis and transpiration (46). Air temperature drives VPD (46), but also affects leaf metabolism, including the ratio of photosynthesis to photorespiration (47). Precipitation is a proxy for soil moisture content and, thus, the amount of water available for uptake by plant roots. Thus, the individual effects of  $C_a$ ,  $VPD_{grw}$ ,  $PPT_{grw}$ , and  $TMP_{grw}$  on our observations of iWUE over the last 115 y are grounded in well-established plant physiology (6, 7, 48, 49). Here, we show that these environmental drivers interact to regulate tree iWUE and are dependent on leaf and wood functional types, demonstrating the complexity and nuance of tree responses to environmental change. In many instances, the effects of increasing  $C_a$  were modulated either positively or negatively by the other drivers of iWUE and by leaf and wood functional types. There was a  $C_a$  by  $VPD_{grw}$  interaction across the 115-y chronology (*SI Appendix, Fig. S6A*) and when considering only the chronology after the 1963 breakpoint (Fig. 24), whereby  $VPD_{grw}$  had a greater effect on iWUE at lower  $C_a$  than at higher  $C_a$  and the effect of increasing  $C_a$  on iWUE was diminished by greater  $VPD_{grw}$ . This may, in part, explain the diminishing influence of  $C_a$  on iWUE presented by Adams et al. (32) in recent years if  $VPD_{grw}$  has also been increasing at the study locations. In contrast, the effect of  $C_a$

**Table 1. LME model results and parameters for the best model (lowest AICc) examining the drivers of tree ring-derived iWUE for the period 1963–2015**

Parameter	1963–2015	1963–2015 <sub>wood</sub>	1963–2015 <sub>leaf</sub>
Intercept, $\mu\text{mol}\cdot\text{mol}^{-1}$	83.078 ± 1.401***	87.297 ± 1.292***	87.592 ± 1.326***
$C_a$ , ppm	0.238 ± 0.007***	0.238 ± 0.007***	0.238 ± 0.007***
$PPT_{grw}$ , mm	-0.010 ± 0.001***	-0.010 ± 0.001***	-0.010 ± 0.001***
$TMP_{grw}$ , °C	0.246 ± 0.135 <sup>ns</sup>	0.245 ± 0.135 <sup>ns</sup>	0.245 ± 0.135 <sup>ns</sup>
$VPD_{grw}$ , kPa	13.141 ± 1.476***	13.131 ± 1.477***	13.131 ± 1.477***
$CO_2$ : $TMP_{grw}$	0.012 ± 0.007 <sup>ns</sup>	0.012 ± 0.007 <sup>ns</sup>	0.012 ± 0.007 <sup>ns</sup>
$CO_2$ : $VPD_{grw}$	-0.175 ± 0.073*	-0.172 ± 0.073*	-0.172 ± 0.073*
$TMP_{grw}$ : $VPD_{grw}$	1.919 ± 0.915*	—	1.936 ± 0.915*
Diffuse porous	—	-17.641 ± 2.851***	—
Ring porous	—	-18.321 ± 2.727***	—
Needleleaf deciduous	—	—	-3.064 ± 2.667 <sup>ns</sup>
Broadleaf evergreen	—	—	-10.720 ± 7.954 <sup>ns</sup>
Broadleaf deciduous	—	—	-18.952 ± 2.285***
Marginal $R^2$	0.12	0.38	0.39
Conditional $R^2$	0.88	0.89	0.89

LME model results and parameters for the best model (lowest AICc) examining the drivers of tree ring-derived iWUE for the period 1963–2015 when only considering environmental factors, climate, and  $CO_2$  (second column), including wood type as a fixed effect with no interactions in the model (third column), and when including leaf type as a fixed effect with no interactions in the model (fourth column). Intercept values for each model represent the value of iWUE when each numerical environmental factor included in the model is at its mean value during the study period. Leaf and wood type parameter estimates indicate the difference from the original intercept value for each respective variable, with comparisons made with “conifers” at the base level for wood type and “needleleaf evergreen” as the base level for leaf type. The marginal  $R^2$  describes the goodness of model fit given fixed effects only, while the conditional  $R^2$  describes the goodness of model fit including fixed and random effects (tree species nested within site). Model parameter significance is denoted by an asterisk, where \* $P < 0.05$ , \*\* $P < 0.01$ , and \*\*\* $P < 0.001$ , and “ns” denotes not significant.

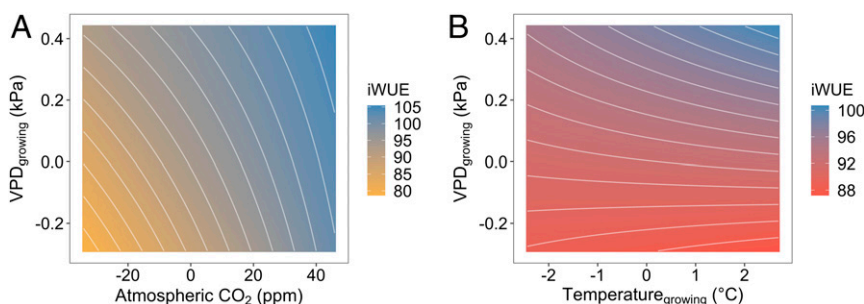
on iWUE was enhanced at greater temperatures across the 115-y chronology (SI Appendix, Fig. S6B).

Furthermore, we found increasing  $C_a$  led to greater increases in iWUE in conifer trees, with respect to trees that had diffuse porous or ring porous wood, and the effect of  $C_a$  was diminished in broadleaf deciduous trees relative to needleleaf evergreen trees (SI Appendix, Table S3). When considering leaf type, we found that  $C_a$  interacted with  $PPT_{grw}$  where the effects of increasing  $C_a$  on iWUE were greatest with low  $PPT_{grw}$  and diminished at high  $PPT_{grw}$  for needleleaf evergreen (Fig. 3A) and needleleaf deciduous (Fig. 3B), in contrast to broadleaf deciduous trees where the effect of increasing  $C_a$  offset the negative effects of high  $PPT_{grw}$  on iWUE (Fig. 3C). These results suggest that iWUE of broadleaf deciduous trees may remain high as  $C_a$  continues to increase regardless of  $PPT_{grw}$ , whereas high  $PPT_{grw}$  offsets the  $C_a$  response of iWUE in conifers, which are adapted to drier environments (50). We found no interaction between  $C_a$  and  $PPT_{grw}$  on iWUE of broadleaf evergreen trees, which instead showed an interaction between  $C_a$  and  $VPD_{grw}$  on iWUE (Fig. 3D). In other instances, the effects of air temperature were modulated by either  $VPD_{grw}$  or precipitation. The response of tree iWUE to  $TMP_{grw}$  after 1963 was dependent upon  $VPD_{grw}$  (Table 1), such that the effect of  $TMP_{grw}$  on iWUE was the greatest at high  $VPD_{grw}$  (Fig. 2B). We found more complex interactions between  $TMP_{grw}$  and  $PPT_{grw}$  on iWUE when considering wood functional types, whereby the greatest iWUE for conifer tree species was at low  $PPT_{grw}$  and low  $TMP_{grw}$  (SI Appendix, Fig. S7A), but at low  $PPT_{grw}$  and high  $TMP_{grw}$  for diffuse porous trees (SI Appendix, Fig. S7B). On the other hand, ring porous trees showed an interaction between  $VPD_{grw}$  and  $PPT_{grw}$ , where iWUE was greatest at low  $PPT_{grw}$  and high  $VPD_{grw}$  (SI Appendix, Fig. S7C). Thus, these interactive effects are important to understand as they underscore the intricate interplay between  $C_a$ , temperature, precipitation, and evaporative water demand on tree physiology, highlight important differences in tree responses to environmental change across tree functional traits, and have important implications for tree carbon gain and water loss as climate changes and  $C_a$  continues to rise (46, 49).

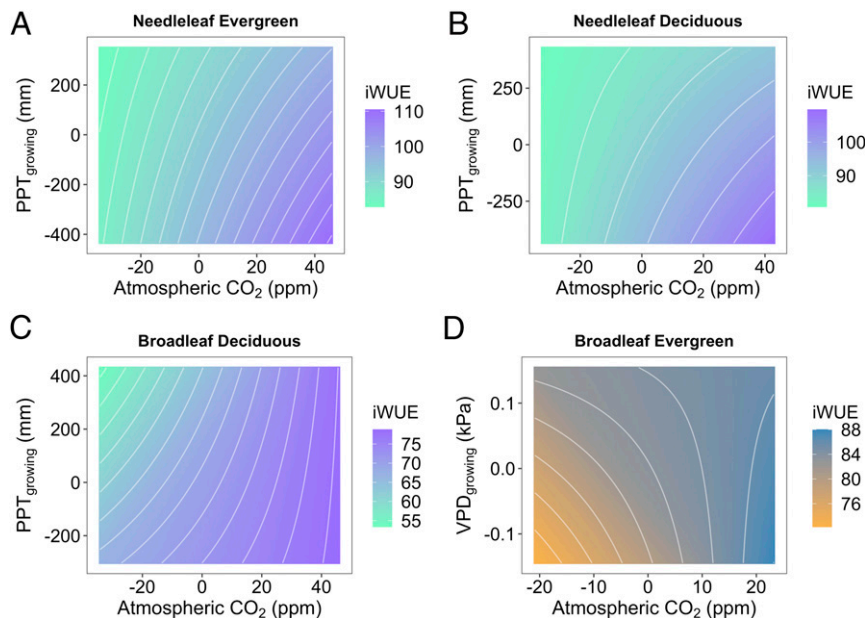
Annually resolved, canopy-integrated iWUE chronologies reconstructed from tree ring carbon isotope signatures clearly show a positive trend over the 20th century explained, in part, by interactions between  $C_a$ ,  $PPT_{grw}$ ,  $VPD_{grw}$ , and  $TMP_{grw}$ . However, whether the increases in iWUE were due to underlying stimulated  $A_{net}$ , reduced  $g_s$ , or some combination thereof, cannot be explained using carbon isotope signatures alone. Therefore, we combined analyses of the chronologies of iWUE with those of  $\Delta^{18}O_{lw}$  (which is inversely related to  $g_s$ ) to partition the increases in iWUE between independent changes in  $A_{net}$  and  $g_s$ . Of all chronologies showing an increase in iWUE since 1963 (Fig. 4A–C), we found 5.3% ( $n = 6$ ) of the studies showed decreasing  $\Delta^{18}O_{lw}$  (Fig. 4D), indicating increased  $g_s$  across the 53-y period; 77.9% ( $n = 88$ ) showed constant

$\Delta^{18}O_{lw}$  (Fig. 4E), indicating no significant change in  $g_s$ ; and 16.8% ( $n = 19$ ) showed increasing  $\Delta^{18}O_{lw}$  (Fig. 4F), reflecting a decrease in  $g_s$  (22, 23, 25, 30, 51). Trends in  $\Delta^{18}O_{lw}$  over time were not different among wood types ( $F = 1.39$ ,  $P = 0.25$ ) (SI Appendix, Fig. S8A), although there were differences due to leaf type ( $F = 4.60$ ,  $P = 0.03$ ), with broadleaf deciduous trees showing slightly increasing  $\Delta^{18}O_{lw}$ , corresponding to a decrease in  $g_s$ , but all other leaf types being no different from each other and having either a nonsignificant slope or a negative slope (SI Appendix, Fig. S8B). There were no relationships between individual trends in  $\Delta^{18}O_{lw}$  and mean  $TMP_{grw}$  and mean  $PPT_{grw}$  across all chronologies since 1963 (SI Appendix, Fig. S9A and B), nor was there a relationship between individual trends in  $\Delta^{18}O_{lw}$  and individual trends in  $TMP_{grw}$  (SI Appendix, Fig. S9D). Mean  $\Delta^{18}O_{lw}$ , however, did decrease with increasing mean  $VPD_{grw}$ , although this relationship was overwhelmingly driven by a few sites ( $n = 5$ ) (SI Appendix, Fig. S9C). Finally, trends in  $\Delta^{18}O_{lw}$  tended to become more negative, corresponding to increased  $g_s$ , in sites becoming wetter since 1963 (SI Appendix, Fig. S9E) and more positive, corresponding to reduced  $g_s$ , in sites where VPD was increasing (SI Appendix, Fig. S9F), although the variability in  $\Delta^{18}O_{lw}$  trends explained by  $PPT_{grw}$  trends and  $VPD_{grw}$  trends was only 9% and 8%, respectively. Analysis of the annual variability in  $\Delta^{18}O_{lw}$  for each of the 113 chronologies since 1963 showed a positive relationship with  $VPD_{grw}$  (SI Appendix, Fig. S10) and  $TMP_{grw}$  (SI Appendix, Fig. S11), but a negative relationship with  $PPT_{grw}$  (SI Appendix, Fig. S12). Our findings are similar to Guerrieri et al. (9) who showed negative or constant  $\Delta^{18}O_{lw}$  trends in wetter sites (increased or constant  $g_s$ , respectively) and align in magnitude and direction with the relationship between VPD and tree ring  $\delta^{18}O$  presented for tropical trees by Kahmen et al. (52). Our analysis of  $\Delta^{18}O_{lw}$  assumed 1) the oxygen isotopic composition of tree source water reflects that of precipitation, and 2) 40% of oxygen atoms exchange with stem water during cellulose synthesis ( $p_{ex} = 0.40$ ) (53). We performed two sensitivity analyses to explore these assumptions, first by allowing a partial decoupling of precipitation and source water oxygen isotope composition, and second by changing  $p_{ex}$  across a range from 0.20 to 0.60, or by a climate-dependent value using the equation provided by Cheesman and Cernusak (54) derived from eucalypts in Northeast Tasmania and found that our overall conclusions using  $p_{ex} = 0.4$  are highly robust (Methods and SI Appendix, Figs. S13–S17 and Table S5).

Whether the increases in tree iWUE are caused by increases in  $A_{net}$ , reductions in  $g_s$ , or a combination of changes in  $A_{net}$  or  $g_s$  are consequential as changes in  $A_{net}$  may affect the carbon cycle possibly through tree growth and carbon sequestration, while changes in  $g_s$  may affect the hydrologic cycle through changes in evapotranspiration. Our temporal analysis of trends in  $g_s$  inferred from individual  $\Delta^{18}O_{lw}$  chronologies from 1963 to 2015 indicates that  $g_s$  increased in 5.3% of examined cases (Fig. 4D) and remained constant in 77.9% (Fig. 4E). Thus, it is necessary that



**Fig. 2.** Nature of the interaction between growing season vapor pressure deficit (VPD) and atmospheric CO<sub>2</sub> (A) and growing season VPD and growing season temperature (B) on iWUE during the period 1963–2015. The interactions shown represent a given predicted value of iWUE throughout the range experienced by each group mean-centered environmental factor since 1963. Values listed for each environmental factor are standardized with respect to the mean during the study period. Parameter estimates for each interaction are listed in Table 1.

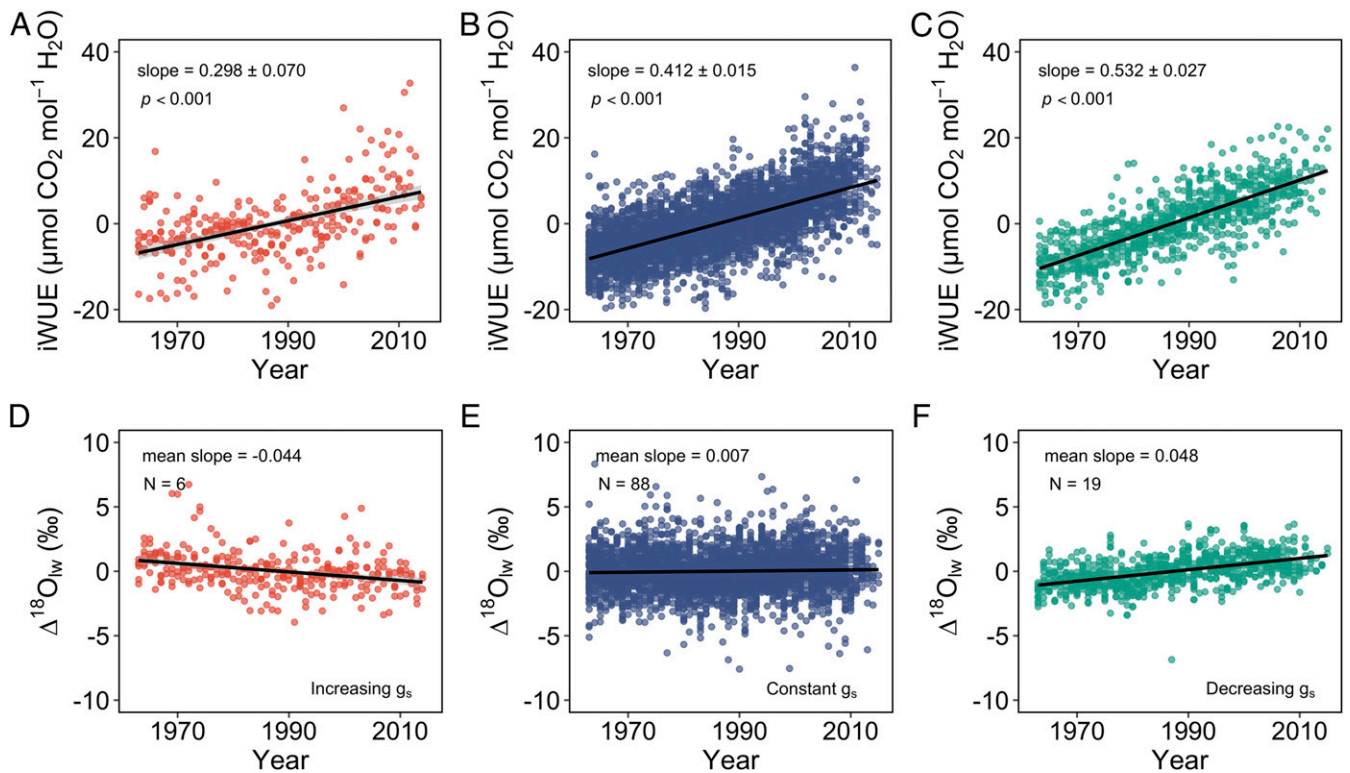


**Fig. 3.** Nature of the interaction between atmospheric CO<sub>2</sub> and growing season precipitation (PPT<sub>growing</sub>) on iWUE for needleleaf evergreen (A), needleleaf deciduous (B), and broadleaf deciduous (C) tree species, and between atmospheric CO<sub>2</sub> and growing season VPD (VPD<sub>grw</sub>) for broadleaf evergreen trees (D). The interactions shown represent a given predicted value of iWUE throughout the range experienced by each group mean-centered environmental factor since 1963. Values listed for each environmental factor are standardized with respect to the mean during the study period.

$A_{net}$  had to increase in 83.2% of examined chronologies (i.e., sum of 5.3% and 77.9%) for iWUE to increase over the 53-y period. In the remaining 16.8% of the chronologies,  $g_s$  declined (Fig. 4F), and as such, increases in iWUE could have occurred because of reductions in  $g_s$  alone or in combination with increasing  $A_{net}$  (Fig. 4C and F). The widespread historical increase in  $A_{net}$  driving increases in tree iWUE was a surprising result to us since declines in  $g_s$ , and a consequent increase in iWUE, are common results found in elevated  $C_a$  experiments (14, 15, 55, 56). However, elevated  $C_a$  studies often rely on large step increases in  $C_a$ , whereas tree rings record responses to long-term progressive, but small, increases in  $C_a$ . Furthermore, analyses of  $g_s$  responses to increased  $C_a$  in trees indicate a large amount of variability with older trees showing less sensitivity than younger trees and conifers being less sensitive than deciduous trees (55). Indeed, ~70% of the chronologies in this study that showed decreasing  $g_s$  were broadleaf deciduous tree species. Our data using tree ring isotopes provide strong support of studies using carbonyl sulfide (13, 42, 57), satellite data (58), and seasonal  $C_a$  patterns (59) that show global increases in  $A_{net}$  as a result of increasing  $C_a$  (13, 42), and build upon recent observations showing widespread stimulated  $A_{net}$  resulting in increased iWUE in the United States (9). Moreover, the rates at which iWUE increased were highest in those chronologies with reduced  $g_s$  (Fig. 4C), followed by those with constant  $g_s$  (Fig. 4B), and the lowest in those chronologies with increased  $g_s$  (Fig. 4A). This highlights the importance of stimulated  $A_{net}$  in driving increasing iWUE in all cases, and supports reductions in  $g_s$ , inferred through increasing  $\Delta^{18}O_{lw}$ , exacerbating realized increases in iWUE in a small subset of chronologies. Finally, for those tree ring chronologies where no trends in iWUE were observed since 1963 ( $n = 18$ ), the overwhelming majority of associated  $\Delta^{18}O_{lw}$  chronologies were either constant ( $n = 16$ ) or decreased ( $n = 1$ ), suggesting any increases in  $A_{net}$  that may have occurred over the 53-y time period were not sufficient to offset the increase or constant  $g_s$ .

Our meta-analysis using historical  $\Delta^{13}C$  and  $\Delta^{18}O_{lw}$  from tree ring chronologies representing 34 species across 10 biomes establishes a strong process-based framework for predicting changes to tree iWUE across biomes and across wood and leaf functional types. It

further provides an extensive annual record documenting a ~40% increase in tree iWUE globally over the 20th century (Fig. 1), similar in magnitude to the ~34% increase in  $C_a$  that occurred over the same time (31). The rate at which iWUE increased more than tripled after 1963 and occurred within years of a similar breakpoint in  $C_a$  (SI Appendix, Fig. S2). Generally,  $C_a$  stimulated the rate of increased iWUE of conifers, which comprised ~74% of examined chronologies, to a greater degree than trees with other wood anatomy. We identified increasing  $C_a$  as a main factor in driving increases in iWUE, although both  $C_a$  and  $TMP_{grw}$  interacted with other environmental drivers of iWUE in ways that suggest trees in areas that experience future increases in PPT<sub>grw</sub> or reductions VPD<sub>grw</sub> may have lower realized iWUE than those areas experiencing drier conditions or higher evaporative demand (Table 1). Metadata on leaf area index (6), tree level photosynthesis or hydraulic conductivity with age (60, 61), height or size effects on carbon isotope discrimination (62), or the levels of air pollution (19, 24, 63) were not consistently available from the published studies used in our analyses, and thus we cannot exclude potential effects of these factors on historical tree iWUE variance. Coupling iWUE with  $\Delta^{18}O_{lw}$  chronologies revealed increases in  $A_{net}$  as a consistent driver behind increasing iWUE across ~83% of the examined chronologies, while reduced  $g_s$ , or a combination of increasing  $A_{net}$  and reduced  $g_s$ , was responsible for increases in iWUE in the remaining ~17% (Fig. 4). The widespread patterns of stimulated  $A_{net}$  from this study are in line with recent findings of a 31% increase in global photosynthetic carbon gain over the 20th century (42), directly tracking increasing  $C_a$  (13, 42). Thus, this meta-analysis showing increased iWUE over the 20th century encompassing a spectrum of tree functional types across a broad geographic area highlights the complexity of tree responses to environmental change and reinforces the importance of stimulated photosynthesis, and not reductions in leaf  $g_s$ , as the primary driver in global increases in iWUE. These results provide a historical baseline of the environmental drivers and physiological mechanisms that result in the uptake of ~30% of anthropogenic carbon emissions by terrestrial ecosystems each year (64).



**Fig. 4.** Standardized chronologies of iWUE (A–C) and  $\Delta^{18}\text{O}_{\text{lw}}$  (D–F) binned by the individual  $\Delta^{18}\text{O}_{\text{lw}}$  trend for each chronology for the period 1963–2015. Red data points (A and D) contain individual chronologies with decreasing  $\Delta^{18}\text{O}_{\text{lw}}$ , blue data points (B and E) contain individual chronologies with constant  $\Delta^{18}\text{O}_{\text{lw}}$ , and green data points (C and F) contain individual chronologies with increasing  $\Delta^{18}\text{O}_{\text{lw}}$ . The respective slope for iWUE across all chronologies within each  $\Delta^{18}\text{O}_{\text{lw}}$  category (decreasing, constant, increasing  $\Delta^{18}\text{O}_{\text{lw}}$ ) is listed within each respective panel, along with the  $P$  value with the corresponding LME model fit where species chronology is nested within site as a random factor (see *SI Appendix, Table S1* for site chronologies). The mean slope across all  $\Delta^{18}\text{O}_{\text{lw}}$  chronologies ( $N$ ) within a given category is listed in the *Top Left* of each panel (D–F). The solid black line in each panel represents the average trend across all data points.

## Materials and Methods

**Data Collection and Extraction.** Data were collected from published articles resulting from a literature search on Web of Science using the search phrase “tree ring, carbon, oxygen, isotope.” Articles published prior to the search date (July 15, 2019) were identified for potential use in our global analysis, and of the 378 candidate articles resulting from the search, 49 met all selection criteria for inclusion, which were as follows: 1) present raw, uncorrected  $\delta^{13}\text{C}$  or  $\delta^{13}\text{C}$ -derived variables (i.e.,  $\Delta^{13}\text{C}$ , iWUE) and  $\delta^{18}\text{O}$  stable isotope signatures from absolutely dated tree ring chronologies, 2) provide identification information regarding tree species sampled and the respective tree ages, 3) include geographic coordinates of each study location, 4) contain high-quality figures from which the isotope data could be extracted or be accompanied by a publicly available supplemental data file, 5) contain metadata regarding the type of tree ring sample used (i.e.,  $\alpha$ -cellulose, bulk wood), and 6) be sampled from mature trees (>50 y) that were not subject to experimental treatments. Our literature search yielded 113 tree ring carbon and oxygen isotope chronologies comprising 36 tree species at 84 sites across the globe (*SI Appendix, Table S1*). Of these 113 chronologies, 84 were conifer tree species and 29 were deciduous tree species, and all but 3 chronologies were from the Northern Hemisphere.

We used WebPlotDigitizer, a data extraction application with built-in tools to extract data from a variety of plot types, with the ability to save files in TAR or JSON format (65), to retrieve  $\delta^{13}\text{C}$ -derived and  $\delta^{18}\text{O}$  isotope data from published figures. We first validated the accuracy and precision of extracted data using figures from our own previously published manuscripts (linear regression,  $R^2 = 1.0$ ; slope = 1.0004;  $y$ -intercept = 0.001;  $P < 0.001$ ). Then, from each article that met the selection criteria, we captured a high-quality image file for each relevant figure in TIFF format and assigned numerical values to the  $x$  axis and  $y$  axis, respectively. This was accomplished by assigning values to the known endpoints of each axis, in turn, and linearly interpolating values between the points. We then identified each unique data point on a given figure, after which the abscissa and ordinate were

automatically calculated. This was performed for each tree- or species-specific chronology within a given site. Typically, species-averaged chronologies for  $\delta^{13}\text{C}$ -derived and  $\delta^{18}\text{O}$  isotope data of all sampled trees were presented in published literature; however, in cases where individual tree chronologies were presented, we first extracted data for each year at the tree level before calculating a species-level mean for a given species within a given site, for direct comparison among observations across all studies.

Environmental factors examined as potential drivers of tree physiology in this study include atmospheric  $\text{CO}_2$  ( $C_a$ , in parts per million), as well as climate factors temperature (TMP, in degrees Celsius), precipitation (PPT, in millimeters), and VPD (in kilopascals) given their relevance to large-scale controls over the carbon and hydrologic cycles (13, 49, 66, 67). Annual  $C_a$  concentrations were obtained from Scripps Institution of Oceanography ice-core merged data product (31, 68), while climate data were obtained from the Climatic Research Unit (CRU TS4.03) global gridded data product (69). Climate data from the CRU TS4.03 data product were provided at a  $0.5^\circ \times 0.5^\circ$  spatial resolution at monthly time step from 1901 to 2018 and were validated using local observations from ground-based meteorological stations. From the CRU TS4.03, we extracted TMP, PPT, and vapor pressure (VAP) (in hectopascals) for each of the 84 unique study locations, and calculated VPD from VAP and TMP using methods from Allen (70). We calculated annual ( $X_{\text{ann}}$ ) and growing season ( $X_{\text{grw}}$ ) values for each climate variable over the period 1901–2015 using means for TMP and VPD, while using sums for PPT. The range of months over which the growing season occurs is dependent upon the geographical site location and was determined based on information provided in each published article.

Last, we compiled characteristics unique to each site and chronology included in this analysis such as biome type, tree ring type, and tree phenology type (*SI Appendix, Table S1*). Biome type was determined based on GPS coordinates provided for each unique site coupled with a global geographical biome classification database outlined in Dinerstein et al. (71),

while tree ring type was determined based upon known ring anatomy confirmed using The Xylem Database (72).

**Carbon Isotope Calculations.** The extracted  $\delta^{13}\text{C}$ -derived chronologies fell into four distinct categories—raw tree ring  $\delta^{13}\text{C}$  ( $\delta^{13}\text{C}_{\text{raw}}$ ,  $n = 45$ ), atmospheric  $\text{CO}_2$ -corrected  $\delta^{13}\text{C}$  ( $\delta^{13}\text{C}_{\text{cor}}$ ,  $n = 36$ ), carbon isotope discrimination ( $\Delta^{13}\text{C}$ ,  $n = 25$ ), and  $\text{iWUE}$  ( $n = 7$ ). While there is a consistent postphotosynthetic fractionation ( $\sim 2\%$ ) between leaf organic matter and carbon assimilated into woody biomass (6, 73, 74), in addition to fractionation related to photorespiration (75), none of the published studies identified in our analysis took this into account when presenting values of  $\delta^{13}\text{C}_{\text{cor}}$ ,  $\Delta^{13}\text{C}$ , or  $\text{iWUE}$ . As such, all published values of  $\delta^{13}\text{C}_{\text{cor}}$ ,  $\Delta^{13}\text{C}$ , and  $\text{iWUE}$  were first converted to  $\delta^{13}\text{C}_{\text{raw}}$  before calculation of  $\delta^{13}\text{C}$ -derived chronologies that included these effects following Lavergne et al. (76). We calculated  $\delta^{13}\text{C}_{\text{raw}}$  when values of atmospheric  $\delta^{13}\text{C}$ -corrected  $\delta^{13}\text{C}$  were given using the following:

$$\delta^{13}\text{C}_{\text{raw}} = \delta^{13}\text{C}_{\text{cor}} + (\delta^{13}\text{C}_{\text{atm}} + 6.4), \quad [1]$$

and when discrimination,  $\Delta^{13}\text{C}$ , was presented using the following:

$$\delta^{13}\text{C}_{\text{raw}} = \frac{1,000 * \Delta^{13}\text{C} - 1,000 * \delta^{13}\text{C}_{\text{atm}}}{1,000 + \Delta^{13}\text{C}}, \quad [2]$$

where  $\delta^{13}\text{C}_{\text{atm}}$  is the carbon isotopic signature of atmospheric  $\text{CO}_2$  on the year of ring formation ( $-6.4\%$  being the  $\delta^{13}\text{C}_{\text{atm}}$  signature prior to the industrial revolution) (77, 78). For all calculations during the period 1901–2003, we used McCarroll and Loader (78) for  $\delta^{13}\text{C}_{\text{atm}}$  values. Beyond 2003,  $\delta^{13}\text{C}_{\text{atm}}$  values were calculated linearly extrapolating the trend beginning at the 1962  $\delta^{13}\text{C}_{\text{atm}}$  breakpoint. In cases where  $\text{iWUE}$  was presented, we first calculated  $\Delta^{13}\text{C}$  using the following:

$$\Delta^{13}\text{C} = \frac{\frac{-\text{iWUE}}{0.625} * (b - a) + C_a * (b - a)}{C_a}, \quad [3]$$

where  $a$  (4.4‰) is the fractionation associated with the diffusion of  $\text{CO}_2$  (77, 79),  $b$  (28‰) is the mean fractionation related to the preferential utilization of  $^{12}\text{CO}_2$  over  $^{13}\text{CO}_2$  during carboxylation by Rubisco (75, 77), and  $C_a$  is the atmospheric  $\text{CO}_2$  concentration during the year of ring formation (31, 68). We then used Eq. 2 to calculate  $\delta^{13}\text{C}_{\text{raw}}$  from  $\Delta^{13}\text{C}$  (77, 80). To calculate  $\text{iWUE}$  from  $\delta^{13}\text{C}_{\text{raw}}$  including photorespiration and postphotosynthetic fractionation effects, we first calculated leaf internal  $\text{CO}_2$  ( $C_i$ ) using the following:

$$C_i = \frac{\frac{(\delta^{13}\text{C}_{\text{atm}} - (\delta^{13}\text{C}_p - d))}{\left(1 + \frac{(\delta^{13}\text{C}_p - d)}{1,000}\right)} - a - f * \left(\frac{\Gamma^*}{pC_a}\right)}{b - a} * C_a, \quad [4]$$

where  $\delta^{13}\text{C}_p$  is the isotopic signature of the tree ring,  $d$  accounts for the carbon isotope fractionation from leaf to wood (bulk wood = 1.9‰; cellulose = 2.1‰) (6, 73),  $f$  is the isotopic fractionation associated with photorespiration (12‰) (75),  $pC_a$  is the partial pressure of atmospheric  $\text{CO}_2$  (in pascals), and  $\Gamma^*$  is the  $\text{CO}_2$  compensation point in the absence of dark respiration (in pascals) (76). We calculated  $\Gamma^*$  according to the following:

$$\Gamma^* = \Gamma_{25}^* * \left(\frac{P_{\text{atm}}}{P_0}\right) * e^{\left(\frac{\Delta H_a * (T - 298)}{R * T * 298}\right)}, \quad [5]$$

where  $\Delta H_a$  is the energy of activation (37,830  $\text{J} \cdot \text{mol}^{-1}$ ) (81),  $T$  is the leaf temperature (in degrees Celsius),  $R$  is the molar gas constant (8.314  $\text{J} \cdot \text{mol}^{-1} \cdot \text{K}$ ),  $P_{\text{atm}}$  is the ambient atmospheric pressure (in pascals),  $P_0$  is the atmospheric pressure at sea level (101,325 Pa), and  $\Gamma_{25}^*$  is the  $\text{CO}_2$  compensation point at 25 °C (4.332 Pa) (81, 82). We assumed  $T$  reflects the mean air temperature over the growing season. We then calculated  $\text{iWUE}$  (in micromoles of  $\text{CO}_2 \cdot \text{moles}^{-1} \text{H}_2\text{O}$ ) as follows:

$$\text{iWUE} = \frac{(C_a - C_i)}{1.6}, \quad [6]$$

where 1.6 is the constant ratio for the diffusivity of water vapor and  $\text{CO}_2$  in air. For each of the preceding calculations, we assumed leaf structure, mesophyll conductance ( $g_m$ ), and mesophyll  $\text{CO}_2$  fractionation remained unchanged throughout the length of each of the chronologies (83); however, we do realize minor species differences in leaf morphology or small changes over time due to environmental influences could have minor influences over  $g_m$ , and thus  $\delta^{13}\text{C}$ -derived chronologies (84). Although it could increase uncertainty around our estimates of  $C_i$ , and thus  $\text{iWUE}$ , this assumption is

nevertheless standard across dendroisotopic studies reconstructing  $\text{iWUE}$  from tree ring carbon isotopes (6, 8, 9, 18, 19, 76).

**Oxygen Isotope Calculations.** We calculated tree ring  $^{18}\text{O}$  enrichment above the source water ( $\Delta^{18}\text{O}_{\text{tr}}$ ) from raw tree ring  $\delta^{18}\text{O}$  ( $\delta^{18}\text{O}_{\text{tr}}$ ) using the following:

$$\Delta^{18}\text{O}_{\text{tr}} = \frac{\delta^{18}\text{O}_{\text{tr}} - \delta^{18}\text{O}_p}{1 + \left(\frac{\delta^{18}\text{O}_p}{1,000}\right)}, \quad [7]$$

where  $\delta^{18}\text{O}_p$  is the oxygen isotopic composition of precipitation during ring formation (54). Given the lack of long-term records of  $\delta^{18}\text{O}_p$  at the study locations, we estimated site-specific  $\delta^{18}\text{O}_p$  over the length of each respective chronology following methods from Guerrieri et al. (9) using the following:

$$\delta^{18}\text{O}_p = 0.52 * T_a - 0.006 * T_a^2 + 2.42 * P_a - 1.43 * P_a^2 - 0.046 * \sqrt{E} - 13, \quad [8]$$

where  $T_a$  is the mean annual temperature (in degrees Celsius),  $P_a$  is total annual precipitation (in meters), and  $E$  is the elevation (in meters) of the study location above sea level (85) (*SI Appendix, Fig. S18*). We used linear regression to validate our estimated values of  $\delta^{18}\text{O}_p$  using Eq. 8 against those obtained from isoMAP (86) for the period 1985–2015 ( $R^2 = 0.94$ ,  $P < 0.01$ ) (*SI Appendix, Fig. S19*). Finally, from  $\Delta^{18}\text{O}_{\text{tr}}$  we calculated the enrichment of leaf water  $^{18}\text{O}$  above source water ( $\Delta^{18}\text{O}_{\text{lw}}$ ) using the following:

$$\Delta^{18}\text{O}_{\text{lw}} = \frac{\Delta^{18}\text{O}_{\text{tr}} - \varepsilon_{\text{wc}}}{1 - p_x * p_{\text{ex}}}, \quad [9]$$

where  $\varepsilon_{\text{wc}}$  is the temperature-dependent fractionation associated with carbonyl oxygen atoms exchanging with water during wood (i.e., cellulose) synthesis (54, 87) and  $p_x$  is the fraction of stem water at the site of cellulose synthesis, of which a given proportion of oxygen atoms,  $p_{\text{ex}}$ , are exchanged. We calculated  $\varepsilon_{\text{wc}}$  following Sternberg and Ellsworth (53) using the following:

$$\varepsilon_{\text{wc}} = 0.0084 * T_a^2 - 0.51 * T_a + 33.172, \quad [10]$$

while  $p_x \sim 1$ , a value in the range of 0.20 to 0.42, is typically used for  $p_{\text{ex}}$  (27, 53, 88), with higher values of  $p_{\text{ex}}$  often observed in more arid sites (54). For this study, we used a conservative value (0.40) for  $p_{\text{ex}}$  (53), although we acknowledge  $p_{\text{ex}}$  may differ across sites with different climates (54). Thus, we performed two sensitivity analyses to examine how changes in  $p_{\text{ex}}$  subsequently influenced  $\Delta^{18}\text{O}_{\text{lw}}$  described in greater detail below. Trends in calculated values of  $\Delta^{18}\text{O}_{\text{lw}}$  were then used to infer changes in  $g_s$  for each individual chronology based on well-established physiological theory describing the inverse relationship between  $\Delta^{18}\text{O}_{\text{lw}}$  and  $g_s$  (i.e., increasing  $\Delta^{18}\text{O}_{\text{lw}}$  indicates reductions in  $g_s$  and decreasing  $\Delta^{18}\text{O}_{\text{lw}}$  indicates increasing  $g_s$ , while constant  $\Delta^{18}\text{O}_{\text{lw}}$  indicates no significant change in  $g_s$ ) (26–30).

One caveat when interpreting  $\Delta^{18}\text{O}_{\text{lw}}$  calculated from tree rings is that in some cases the environmental signal can be muted (i.e., leaf and wood oxygen isotope signatures may become uncoupled) (89), potentially obscuring the inferred  $g_s$ , especially if precipitation  $\delta^{18}\text{O}$  is decoupled from source water  $\delta^{18}\text{O}$ . We first addressed this issue by using ordinary least-squares regression to examine the relationship between calculated  $\Delta^{18}\text{O}_{\text{lw}}$  and  $\text{PPT}_{\text{grw}}$ ,  $\text{VPD}_{\text{grw}}$ , and  $\text{TMP}_{\text{grw}}$  for each of the 113 unique chronologies used in this study (*SI Appendix, Figs. S10–S12*), which show a similar relationship between  $\Delta^{18}\text{O}_{\text{lw}}$  and  $\text{VPD}_{\text{grw}}$  with Kahmen et al. (52), who found a strong preservation of environmental signals in tree ring oxygen isotope composition. Furthermore, we conducted a sensitivity analysis to determine the degree to which precipitation  $\delta^{18}\text{O}$  being uncoupled from source water  $\delta^{18}\text{O}$  may influence our estimates of  $\Delta^{18}\text{O}_{\text{lw}}$ , the resulting long-term  $\Delta^{18}\text{O}_{\text{lw}}$  trends, and the consequences for inferred changes in  $A_{\text{net}}$  and  $g_s$ , as  $\text{iWUE}$  increased (*SI Appendix, Fig. S13 and Table S5*). To do this, we used calculated values of  $\delta^{18}\text{O}_p$  from Eq. 8 used in our analyses to generate simulated data for each chronology with a resulting Pearson's product-moment correlation with the actual  $\delta^{18}\text{O}_p$  of 0.90, 0.80, 0.70, and 0.60, respectively. These scenarios, in turn, represent a 10%, 20%, 30%, and 40% decoupling of  $\delta^{18}\text{O}$  source water from  $\delta^{18}\text{O}_p$ . We then used ANOVA and analysis of covariance (ANCOVA) to test whether uncoupling influenced mean  $\Delta^{18}\text{O}_{\text{lw}}$ , or the rate of change in  $\Delta^{18}\text{O}_{\text{lw}}$  over time. Last, we binned the chronologies by whether they showed increasing, constant, or decreasing  $\Delta^{18}\text{O}_{\text{lw}}$  as a consequence of uncoupling of  $\delta^{18}\text{O}_p$  and source water  $\delta^{18}\text{O}$  to examine the consequences for our main conclusions regarding changes in  $A_{\text{net}}$  and  $g_s$  (*SI Appendix, Table S5*).

In addition to the potential for precipitation  $\delta^{18}\text{O}$  to become uncoupled from source water  $\delta^{18}\text{O}$ , the proportion of oxygen exchanging with local water during cellulose synthesis ( $p_{\text{ex}}$ ) may vary as a function of climate (54). Our second sensitivity analysis was designed to address how changes in  $p_{\text{ex}}$  influence  $\Delta^{18}\text{O}_{\text{lw}}$  and the overall conclusions drawn from our dual isotope analysis (SI Appendix, Table S5). We performed this first by calculating  $\Delta^{18}\text{O}_{\text{lw}}$  as a constant with  $p_{\text{ex}} = 0.2, 0.4, \text{ and } 0.6$ , and second where  $p_{\text{ex}}$  was allowed to vary as a function of mean annual precipitation, assuming the equation provided by Cheesman and Cernusak (54), derived from eucalypts in Northeast Tasmania, is consistent with values of  $p_{\text{ex}}$  of the chronologies for the sites used in our study (SI Appendix, Figs. S14 and S15). We used ANCOVA to test whether trends in  $\Delta^{18}\text{O}_{\text{lw}}$  over time was dependent upon the value of  $p_{\text{ex}}$  used (SI Appendix, Fig. S16). We then used ordinary least-squares regression to examine the strength, and nature, of the relationship between  $\Delta^{18}\text{O}_{\text{lw}}$  and  $\text{VPD}_{\text{grw}}$  for each  $p_{\text{ex}}$  scenario (SI Appendix, Fig. S17). Results for this analysis are described in the figure legends (SI Appendix, Figs. S14–S17), and these data led us to conclude that varying  $p_{\text{ex}}$  from 0.4 by a constant did not significantly affect the number of  $\Delta^{18}\text{O}_{\text{lw}}$  trends that increased, decreased, or did not change over time. Varying  $p_{\text{ex}}$  by trying to account for different ecosystem climates resulted in four additional chronologies showing decreasing or constant  $\Delta^{18}\text{O}_{\text{lw}}$  and reduced the number of chronologies showing increasing  $\Delta^{18}\text{O}_{\text{lw}}$  by 8, but this did not change the overall conclusions of our study (SI Appendix, Table S5).

**Statistical Analyses.** To examine long-term trends in iWUE, we fit a LME model using the R package *nlme* (90) including *year* as the sole fixed effect with *tree species* nested within *site* as random factors. In this analysis, we examined trends in iWUE across the entire chronology for which climate data were available (full: 1901–2015), in addition to two unique periods (pre: 1901–1963; post: 1963–2015) after we identified a breakpoint in the full iWUE chronology (across all sites and species) using the *segmented* package in R (91). We further assessed breakpoints in individual chronologies, with the requirement there was a  $\pm 25$ -y buffer with respect to 1963 (SI Appendix, Fig. S1). Additionally, we used generalized least-squares (GLS) regression to examine trends in isotope-derived response variables (iWUE,  $\Delta^{18}\text{O}_{\text{lw}}$ , etc.) separately for each species within a given site, as well as for environmental factors (PPT, TMP, etc.) at a given site for each time period (i.e., full, pre, post) using the R package *nlme* (90). Each LME and GLS model used to determine temporal trends were fit via restricted maximum likelihood and included a first-order autocorrelation structure to account for temporal autocorrelation in the chronologies.

We further used LME models to identify which environmental factors were most responsible in driving the observed changes in iWUE over the last

115 y using the R package *nlme* (90). We used the *dredge()* function from the *MuMIn* R package (92) to examine all possible model LME model combinations ( $n = 113$ ) that included the environmental factors  $C_a$ ,  $\text{TMP}_{\text{grw}}$ ,  $\text{PPT}_{\text{grw}}$ ,  $\text{VPD}_{\text{grw}}$ , wood type, and leaf type as independent variables, as well as interactions up to order 2. We identified the top candidate model as having the lowest corrected Akaike information criterion (AIC<sub>c</sub>) for subsequent parameter inference. Prior to this analysis continuous environmental factors were grouped by each unique chronology and centered to their mean to prevent issues that arise due to multicollinearity (9, 93). Each LME model examined included a temporal autocorrelation structure of order 1 [i.e., AR(1,0)], as well as species nested within site as a random factor, and were fit via maximum likelihood for comparison of models with different fixed effects. The top candidate model was then refit via restricted maximum likelihood. We used the R package *visreg* (94) to visualize the nature of significant interactions between environmental factors on iWUE over the range experienced by each factor.

As a complement to LME model parameter inference, we examined the independent effects of each environmental factor included as a fixed effect in the final model on iWUE through hierarchical partitioning (HP) using the R package *hier.part* (95). HP does not consider potential interactions between factors, but instead determines the contribution of a given environmental factor to iWUE through an examination of every possible model structure containing a given predictor variable. By doing so, HP avoids enhancing or diminishing the variance explained by any given predictor variable, and circumvents issues related to multicollinearity among predictor variables (96). To do this, we used the *rand.hp()* function, which randomizes the matrices containing response and predictor variables and computes independent effects, which we set at 1,000 repetitions to determine variable contributions to iWUE.

We performed ANCOVA to examine differences in the rate of change of iWUE and  $\Delta^{18}\text{O}_{\text{lw}}$  over time that depended upon leaf type, wood type, or biome, and performed ANOVA to compare iWUE means between leaf type, wood type, and biome. Furthermore, we used the R package *emmeans* (97) to test for pairwise differences among groups from ANCOVA model fits, and used the R package *multcomp* (98) to generate connecting letters reports from Tukey's honestly significant difference post hoc test for multiple comparisons of means from ANOVA model fits.

**Data Availability.** All tree ring isotope data and code supporting this analysis have been deposited on Figshare (<https://doi.org/10.6084/m9.figshare.13541927>).

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